

POPULAR SUMMARY FOR THE ARTICLE

“Saharan Air and Atlantic tropical cyclone suppression from a global modeling perspective.”

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being submitted to *Geophysical Research Letters*

During summer 2006, the NASA African Monsoon Multidisciplinary Analysis (NAMMA) organized a field campaign in Africa called Special Observation Period (SOP-3), in which scientists in the field were involved in a number of surface network and aircraft measurements. One of the scientific goals of the campaign was to understand the nature and causes for tropical cyclogenesis originating out of African Easterly Waves (AEWs, westward propagating atmospheric disturbances sometimes associated with precursors of hurricanes), and the role that the Saharan Air Layer (SAL, a hot and dry air layer advecting large amounts of dust) can play in the formation or suppression of tropical cyclones. During the NAMMA campaign a high-resolution global model, the NASA GEOS-5, was operationally run by the NASA Global Modeling and Assimilation Office (GMAO) in support to the mission. The daily GEOS-5 forecasts were found to be very useful by decision-making scientists in the field as an aid to discriminate between developing and non-developing AEWs and plan the flight tracks.

In the post-event analyses which were performed mostly by the Goddard Laboratory for Atmospheres, two events were highlighted: a non-developing AEW which appeared to have been suppressed by Saharan air, compared to a developing AEW which was the precursor of hurricane Helene. Both events were successfully predicted by the GEOS-5 during the real-time forecasts provided in support to the mission.

In this work it is found that very steep moisture gradients and a strong thermal dipole, with relatively warm air in the mid-troposphere and cool air below, are associated with SAL in both the GEOS-5 forecasts and the NCEP analyses, even at -great distance- from the Sahara.

The presence of these unusual thermodynamic features over the Atlantic Ocean, at several thousands of kilometers from the African coastline, is suggestive that SAL mixing is very minimal and that the model's capability of retaining the different properties of air masses during transport are important to represent effectively the role of dry air intrusions in the tropical circulation.

¹ **Saharan Air and Atlantic tropical cyclone suppression from a
2 global modeling perspective.**

O. Reale,^{1,2} W. K. M. Lau,¹ A. daSilva,³ K.-M. Kim^{1,2}

³ This article investigates the role of the Saharan Air
4 Layer (SAL) in two cases of non-developing and de-
5 veloping systems observed during the Special Observa-
6 tion Period (SOP-3) phase of the 2006 NASA African
7 Monsoon Multidisciplinary Analyses (NAMMA). A high-
8 resolution global model, the NASA GEOS-5 was opera-
9 tionally run by the NASA Global Modeling and Assim-
10 ilation Office (GMAO) in support to the NAMMA field
11 campaign, which included surface network and aircraft
12 measurements. One of the scientific goals of the cam-
13 paign was to understand the nature and causes for trop-
14 ical cyclogenesis out of African Easterly Waves (AEWs).
15 The daily GEOS-5 forecasts were found to be very use-
16 ful by decision-making scientists in the field as an aid
17 to discriminate between developing and non-developing
18 AEWs and plan the flight tracks. In the post-event anal-
19 yses of a non-developing system which appeared to have
20 been suppressed by SAL, it has been found that very
21 steep moisture gradients and a strong thermal dipole are
22 associated with SAL in both the GEOS-5 forecasts and
23 the NCEP analyses, even at great distance from the Sa-
24 hara. This is suggestive that SAL mixing is very minimal
25 and that the model's capability of retaining the different
26 properties of air masses during transport is important to
27 represent effectively the role of dry air intrusions in the
28 tropical circulation.

1. Introduction

²⁹ African Easterly Waves (AEWs) have been recognized
30 as prominent weather-producing events of northern trop-
31 ical Africa [e.g Burpee, 1974; Asnani, 2005] and have
32 been extensively studied from observational and model-
33 ing perspectives [e.g Hsieh and Cook, 2005; Kiladis et al.,
34 2006].

³⁵ However, the development of AEWs into tropical de-
36 pressions remains one of the most challenging problem in
37 the prediction and modeling of Atlantic tropical cyclones.
38 The Saharan Air Layer (SAL), a layer of hot dry air rich
39 in dust and produced over the Saharan desert, has been
40 investigated with the aid of Geostationary Operational
41 Environmental Satellite (GOES) by Dunion and Velden
42 [2004] and recognized as a possible, important mechanism

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43 concurring to tropical cyclone suppression.

44 The general problem of tropical cyclogenesis has al-
45 ways been considered from either strictly observational
46 or high-resolution mesoscale points of view, because the
47 lower resolution of global models is deemed to be inad-
48 equate to trigger spontaneous cyclogenesis. However, in
49 recent years a number of global models have reached the
50 resolution of 10-40 km and have started to display some
51 tropical cyclogenesis capability [e.g. Atlas et al., 2005;
52 Shen et al., 2006].

53 The organization of several convective centers into
54 a rotating system still requires cloud-resolving models
55 and can not be reproduced by real-time global numer-
56 ical weather prediction models. However, the process of
57 cyclogenesis or cyclogenesis suppression can be studied
58 from different perspectives, and global models do have
59 the advantage of better capturing the large-scale forcings
60 involved. In particular, a global model can be used to
61 investigate the role of SAL not only on the wave scale,
62 but also from the point of view of the large-scale trans-
63 port from its source region, and can therefore represent
64 the possible modification of thermodynamical properties,
65 as the waves propagate over thousands of kilometers. At
66 the same time, sufficiently high resolution is needed to
67 unveil some of the SAL kinematic features, such as the
68 increasingly narrow structure of the dry air filaments be-
69 ing intruded in a tropical circulation, and the sharpness
70 of the boundaries between Saharan and non-Saharan air.

2. The Model

71 In this work, the effect of Sahara air is investigated
72 using a high-resolution global atmospheric model, the
73 NASA GEOS-5, documented in Bosilovich et al. [2006].
74 The GEOS-5 shares the same dynamical core [Lin, 2004]
75 with the so-called NASA finite-volume General Circula-
76 tion Model (fvGCM) which has demonstrated remarkable
77 capabilities in hurricane forecasting [Atlas et al., 2005;
78 Shen et al., 2006]. The GEOS-5 however contains a new
79 physics developed predominantly by the Global Model-
80 ing and Assimilation Office (GMAO), which is substan-
81 tially different from the previous fvGCM. The version
82 used during NAMMA was run at a horizontal resolution
83 of $0.25^\circ \times 0.33^\circ$ with 72 vertical levels.

3. MAP06 and NAMMA

84 Over the past ten years, the NASA Global Modeling
85 and Assimilation Office, continuing the work previously
86 done in the Data Assimilation Office, has been increasing
87 the resolution of its global model and performing real-
88 time forecasts. During the Atlantic tropical seasons of
89 2005, 2006 and 2007 the current high resolution version of
90 the model was put to a severe test by providing real-time
91 forecasts that could be compared to operational state-
92 of-the-art models. GEOS-5 was made available to the
93 operational team in Africa, as an auxiliary decision sup-
94 port tool in the SOP-3 phase of the NAMMA campaign.
95 The model performed well and the forecasts were found
96 beneficial by the team on the field. After-event analysis
97 has revealed some interesting aspects of the cyclogenetic
98 process as perceived by the model, and on the impact of
99 SAL in the development or suppression of storms.

4. Analysis

100 In this work we focus on two interesting events ob-

101 served during the SOP-3 NAMMA campaign: one non-
102 developing and one developing wave, appearing quite sim-
103 ilar in terms of intensity, vertical shear and other dy-
104 namical forcings, and we investigate their different evo-
105 lution and the different properties of the corresponding
106 SAL intrusions. Figure 1 shows a Hovmöller of 850hPa
107 relative vorticity and total precipitable water from the
108 operational NCEP analyses to emphasize AEWs during
109 the first 15 days of the SOP-3, covering the second half
110 of August. The strongest wave of the period appears in
111 the diagram on 23 August at about $5^{\circ}W$ and undergoes
112 transition on the following day (hereafter W1). An ev-
113 ident sharp strip of dry air with the same propagation
114 speed and amplitude of the wave, clearly associated with
115 a Saharan Air outbreak, can be seen.

116 In Figure 2, 700 hPa specific humidity and 850 hPa
117 flow are shown together to emphasize the interaction be-
118 tween two different levels at the initial time (00z 26 Au-
119 gust, from NCEP analyses) and across 3 times of the
120 forecast (24, 48 and 72 hour forecast, corresponding to
121 verification times of 00z 27, 28 and 29 August respec-
122 tively). The 700hPa level is at the lower part of the
123 SAL, whereas 850 corresponds approximately to the top
124 of the moist lower level and emphasizes the low-level cir-
125 culation. Flows at 700 and 850 hPa are substantially
126 different except around the storm center, where a verti-
127 cally aligned circulation is present from the surface up
128 to almost 500 hPa. Based upon these forecasts, it was
129 correctly suggested to the NAMMA forecasting team on
130 the field that W1 would become a nondeveloping wave
131 in spite of its apparent strength. The team obviously
132 had many other forecasting tools available but the infor-
133 mation provided by the GEOS-5 was correct. Following
134 the path of the dry air at 700 hPa it can be seen that
135 as soon as dry air is advected on the top of the 850 hPa
136 circulation center, the rotating system becomes first elon-
137 gated and then rapidly evolves into an open wave. This
138 is even more evident while analyzing intermediate time-
139 steps (not shown).

140 In Figure 3, a zonal vertical cross-section of specific
141 humidity at $20^{\circ}N$ is extracted from the GEOS-5 24-hour
142 forecast, right across the center of the same circulation
143 which can still be seen in Figure 2 at 00z 27 August. An
144 intriguing feature, namely a sharply defined ‘corridor’ of
145 extremely dry air can be seen down to 800 hPa. Re-
146 markable moisture gradients are present on both sides.
147 From the temperature anomaly (obtained by subtract-
148 ing the zonal mean between $80^{\circ}W$ and 0°), a very well-
149 defined thermal dipole, stronger than any other anomaly
150 in the range of longitudes selected, can be seen in per-
151 fect correspondence to the dry tongue. In particular, a
152 warm anomaly spans between 800hPa and 400hPa, and a
153 cool anomaly between 825 hPa and the surface. In other
154 words, since the cross-section cuts across the SAL intru-
155 sion, it appears that temperature, in the core of the SAL,
156 is approximately $3^{\circ}C$ warmer than the surroundings at
157 the same latitude, in partial agreement with Dunion and
158 Velden, [2004]. The new aspect of this analysis is however
159 that a *negative* value also detected in the moist low-level
160 layer at the base of the column.

161 In Figure 4, the same figure is extracted from the full-
162 resolution NCEP operational analyses in model levels.
163 0.5 corresponds roughly to 500 hPa. The Figure confirms
164 the thermal structure depicted in the GEOS-5 24-hour
165 forecast. Despite small scale differences, the NCEP anal-
166 yses confirm the presence of a very well-defined dry intru-
167 sion at about $35^{\circ} - 40^{\circ}W$. Most remarkable is the pres-
168 ence of the same dipole thermal anomaly seen in Figure 3.

169 A positive value reaching $4^{\circ}C$ in the mid-troposphere,
170 and a corresponding cool anomaly down to $-4^{\circ}C$ in the
171 low moist layer. The anomalies are obtained, as in Figure 3, by simply subtracting the $80^{\circ}W - 0^{\circ}$ mean.
172

173 In contrast to Fig 1, the Hovmöller computed for
174 September (Figure s1) shows a powerful wave towards
175 the end of SOP-3 (hereafter W2) associated to some Sa-
176 haran Air: however, the magnitude of the SAL is not
177 comparable with the case of W1.

178 In Figure 5, the same zonal cross-section is produced
179 across the center of the system W2, which is a precur-
180 sor of Helene: the vertically aligned vorticity column
181 at about $33^{\circ}W$ is the analyzed signature of the Trop-
182 ical Storm, named at 00z 14 September 2006 [Brown,
183 2006]. A weak positive temperature anomaly at about
184 $40 - 45^{\circ}W$ in the lower midtroposphere is associated with
185 the same Saharan air outbreak which can be detected in
186 the Hovmöller in Figure S1. However, two prominent dif-
187 ferences can be seen with respect to the non-developing
188 W1 in Figures 3 and 4: there is no cool anomaly in the
189 lowest levels, and the dry air appears more diluted with
190 less sharp horizontal gradients.

191 The real-time GEOS-5 tropical cyclogenesis forecast
192 for Helene was correct. In Figure S2, the 850 hPa circu-
193 lation shows a clearly defined vortex (at about $24 - 25^{\circ}W$
194 and $10 - 12^{\circ}N$ in the initial conditions) progressing west-
195 ward and then recurving northwestward, being entangled
196 in corresponding high levels of 700 hPa moisture. Based
197 upon this and other information, subjective forecasts on
198 the field considered the possibility of that system to be
199 a developing one, and one flight was successfully planned
200 across it.

201 Post-event model analysis suggests that since vertical
202 shear and all other environmental conditions were very
203 favorable in both W1 and W2 case (not shown), but only
204 the latter underwent development becoming Helene, the
205 only difference appears to be the intensity of the Sa-
206 haran Air intrusion. In the model, the temperature dipole
207 associated to the SAL could be followed at each timestep
208 and can be considered a possible cause of suppression. In
209 the precursor of Helene, the low-level negative anomaly
210 was minimal or absent.

211 While the positive anomaly can be simply attributed
212 to the signature of warm air originated over the Sahara,
213 the cool anomaly in the lower levels does not have any
214 plausible explanation relying on transport only. There
215 is no source of localized cooler temperatures at that lat-
216 itude, away from landmass and in a very homogeneous
217 marine tropical environment. At this time, albeit spec-
218 ulative, a possible explanation is that the low-level cool
219 temperatures are an indirect evidence of dust amount.
220 The thermal effect of Saharan mineral dust is a net reduc-
221 tion of downwave shortwave radiation in the near-surface
222 levels, and a heating in the lower midtroposphere, cor-
223 responding to the core of the SAL. It appears that the
224 high-resolution NCEP, unlike lower resolution analyses,
225 can represent this thermal structure and that the GEOS-
226 5 model initialized by the NCEP analyses could retain
227 it for 24-72 hours advecting it into the circulation and
228 producing a realistic cyclone dissipation.

229 In the case of the GEOS-5 forecast, the nature of the
230 finite-volume dynamics [Lin, 2004] is such that is a par-
231 ticularly suitable tool to generally maintain sharp gradi-
232 ents by minimizing unrealistic diffusion processes. The
233 finite-volume dynamics has been shown to be very ef-
234 ficient in the midlatitudes where localized temperature
235 gradients associated with sharp fronts can be very realis-
236 tically simulated and maintained. This work documents

237 that the same skill can be very useful also in the tropics
238 when dealing with Saharan Air.

5. Concluding Remarks

239 This work documents the contribution of the NASA
240 Global Modeling and Assimilation Office in support to
241 the SOP-3 phase of the NAMMA campaign. From the
242 30 5-day forecasts, one prominent case is extracted, a
243 very strong non-developing wave, and is compared with
244 the wave that becomes the precursor of Hurricane He-
245 lene. GEOS-5 forecasts and NCEP full-resolution analy-
246 ses document the presence of a strong temperature dipole
247 (cooler than the environment below 800hPa, and warmer
248 from 800 to 500hPa) associated with the Saharan air in-
249 trusion. This dipole is advected into the circulation of the
250 wave, suppressing further development. No such dipole is
251 found for the Saharan air intruded in the Helene's precur-
252 sor. The lower tropospheric cooling associated with the
253 strong Saharan air outbreak is suggestive that the high
254 resolution global models and analyses can capture part
255 of the thermal effect consequent to downward shortwave
256 reduction caused by large amounts of Saharan dust. An
257 interactive dust aerosol component, including its direct
258 radiative effects, is being employed within GEOS-5 to
259 further assess and quantify the thermal effects of dust on
260 tropical cyclogenesis and will be used for a future study.

261 **Acknowledgments.** Authors acknowledge support from
262 Ramesh Kakar through the NAMMA Project. Thanks are also
263 due to Michele Rienecker for access to the GEOS-5 forecasts.

References

264 Asnani, G. C. (2005), *Tropical Meteorology*. Publisher: Indian
265 Institute of Tropical Meteorology, Pashan, Pune-411008,
266 India. 3 vols.
267 Atlas, R., O. Reale, B.-W. Shen, S.-J. Lin, J.-D.
268 Chern, W. Putman, T. Lee, K.-S. Yeh, M. Bosilovich,
269 and J. Radakovich (2005), Hurricane forecasting with
270 the high-resolution NASA finite-volume general circula-
271 tion model. *Geophysical Research Letters*, **32**, L03807,
272 doi:10.1029/2004GL021513.
273 Bosilovich, M. G., S.D. Schubert, M. Rienecker, R. Todling,
274 M. Suarez, J. Bacmeister, R. Gelaro, G.-K. Kim, I. Stajner,
275 and J. Chen, (2006), NASA's Modern Era Retrospective-
276 analysis for Research and Applications. U.S. CLIVAR Vari-
277 ations, **4** (2), 5-8.
278 Brown, D. P., (2006), Tropical Cyclone Report. Hurricane He-
279 lene (AL082006). 12-24 September 2006. Available online
280 at
281 http://www.nhc.noaa.gov/pdf/TCR-AL082006_Helene.pdf
282 Burpee, R. W. (1974), Characteristics of north African East-
283 erly waves during the summers of 1968 and 1969. *J. Atmos.*
284 *Sci.*, **31**, 1556-1570.
285 Dunion, J., and C. S. Velden, (2004), The impact of the Saha-
286 ran Air Layer on Atlantic Tropical Cyclone activity. *Bull.*
287 *Am. Meteorol. Soc.*, **85**, 353-365.
288 Hsieh, J.-S., and K. Cook (2005), Generation of African East-
289 erly Wave disturbances: Relationship to the African East-
290 erly Jet. *Mon. Wea. Rev.* **133**, 1311-1327.
291 Kiladis, G. N., C. D. Thorncroft, and N. M. J. Hall (2006),
292 Three-Dimensional Structure and Dynamics of African
293 Easterly Waves. Part I:Observations, *J. Atmos. Sci.*, **63**,
294 2212-2230.
295 Lin, S.-J., 2004: A 'vertically lagrangian' finite-volume dy-
296 namical core for global models, *Mon. Wea. Rev.*, **132**, 2293-
297 2307.
298 Shen, B.-W., R. Atlas, O. Reale, S.-J. Lin, J.-D. Chern, J.
299 Chang, C. Henze, and J.-L. Li (2006), Hurricane forecasts

300 with a global mesoscale-resolving model: Preliminary re-
301 sults with Hurricane Katrina (2005). *Geophysical Research*
302 *Letters*, 33, L13813, doi:10.1029/2006GL026143.

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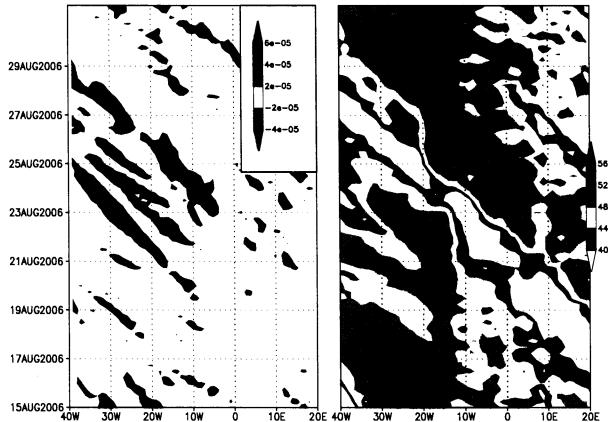


Figure 1. Hovmöller of 850 hPa relative vorticity (s^{-1} , left panel) and total precipitable water (right panel) from the NCEP operational analyses, latitudinally averaged ($10^{\circ} - 18^{\circ} N$), covering the period from 16 to 31 August. Data on pressure levels interpolated on a 1° resolution grid. No significant difference from the original data in sigma levels (not shown).

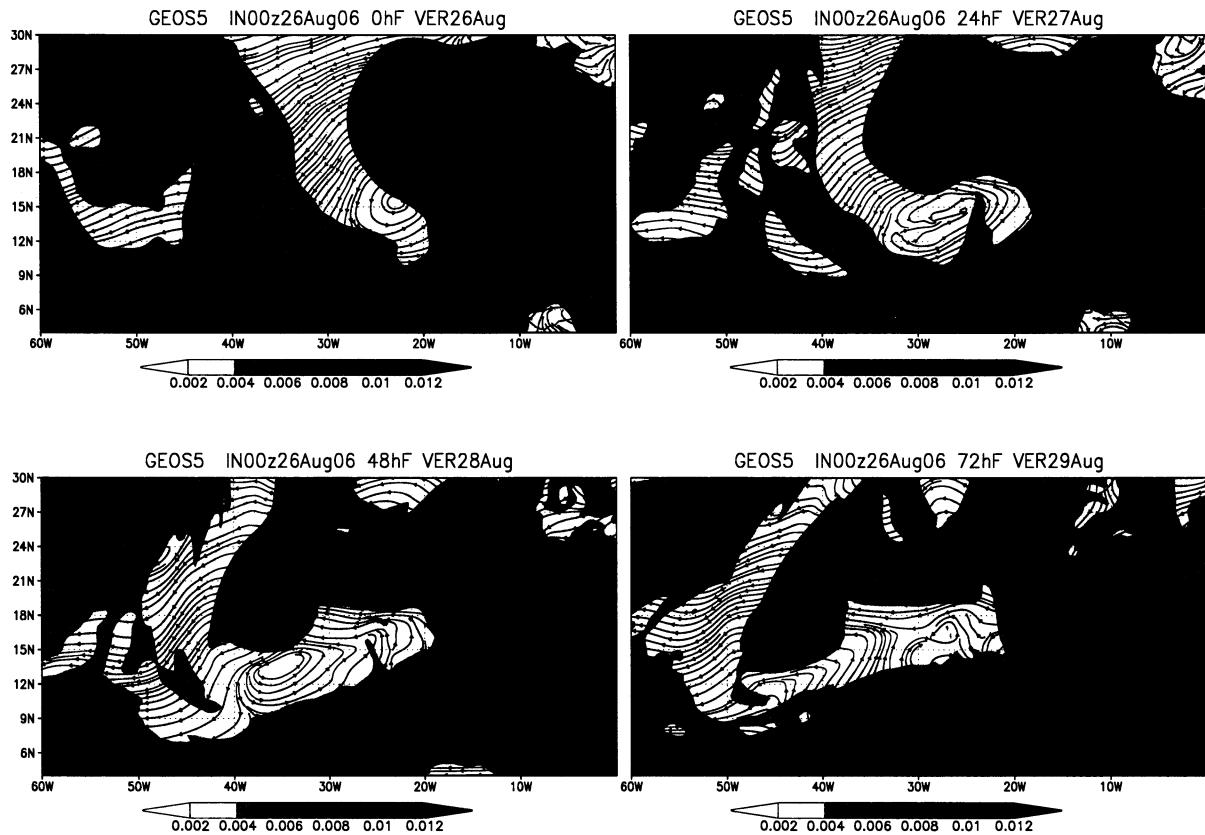


Figure 2. GEOS-5: 700 hPa specific humidity ($KgKg^{-1}$) and 850 hPa wind (streamlines) in the NCEP-derived initial conditions (upper left) for 00z 26 August, and relative to the 24, 48 and 72 hour forecasts for 00z 26, 27 and 28 August.

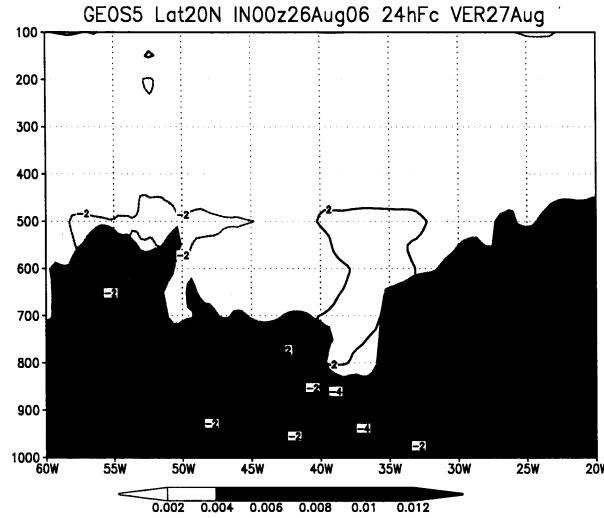


Figure 3. GEOS-5: zonal vertical cross-section of specific humidity ($KgKg^{-1}$, shaded) and temperature anomaly ($^{\circ}C$, contour, subtracting the zonal mean between $80^{\circ}W$ and 0°) at $20^{\circ}N$ for 27 August, 24 hour forecast initialized at 00z 26 August.

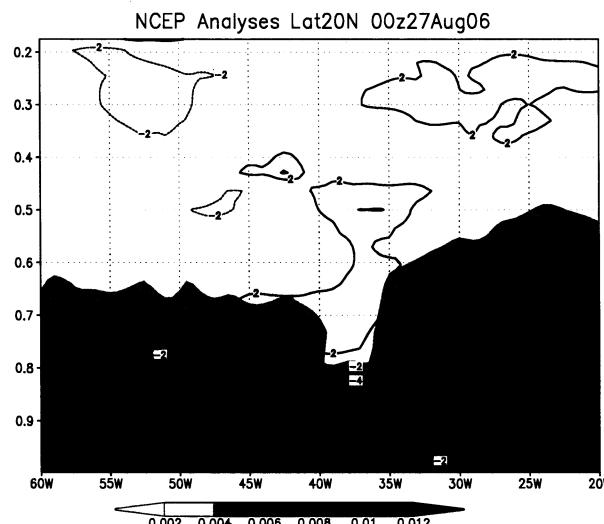


Figure 4. NCEP full-resolution analyses in model levels: zonal vertical cross-section of specific humidity ($KgKg^{-1}$, shaded) and temperature anomaly ($^{\circ}C$, contour, subtracting the zonal mean between $80^{\circ}W$ and 0°) at $20^{\circ}N$ for 00z 27 August. The vertical dimension is only approximately comparable with Figure 3 since the spacing between model levels and pressure levels is different.

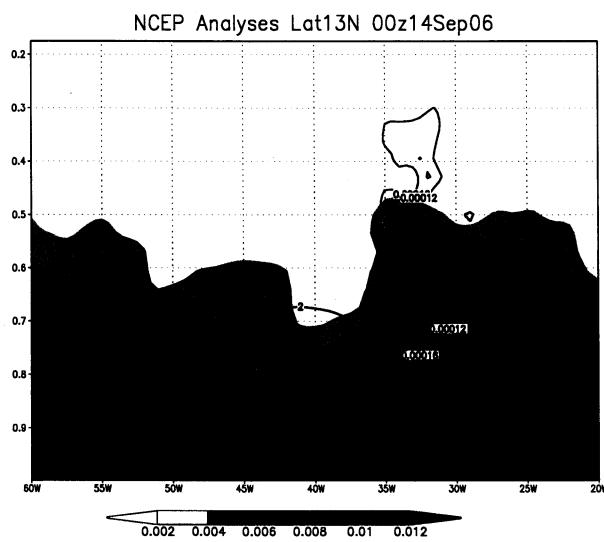


Figure 5. NCEP full-resolution analyses in model levels: zonal vertical cross-section of specific humidity ($Kg\,Kg^{-1}$, shaded), temperature anomaly ($^{\circ}C$, contour, subtracting the zonal mean between $80^{\circ}W$ and 0°) and relative vorticity (s^{-1} , contour) at $13^{\circ}N$ for 00z 14 Sep, across the center of the newly named Tropical Storm Helene.

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Auxiliary Material

Two additional figures are provided to allow a comparison with Fig. 1 and Fig. 5 in the paper.

grl_fs01.eps

Caption:

Hovmoller of 850 hPa relative vorticity (left panel) and total precipitable water (right panel) from the NCEP operational analyses, latitudinally averaged (10-18N), covering the period from 1 to 15 September.

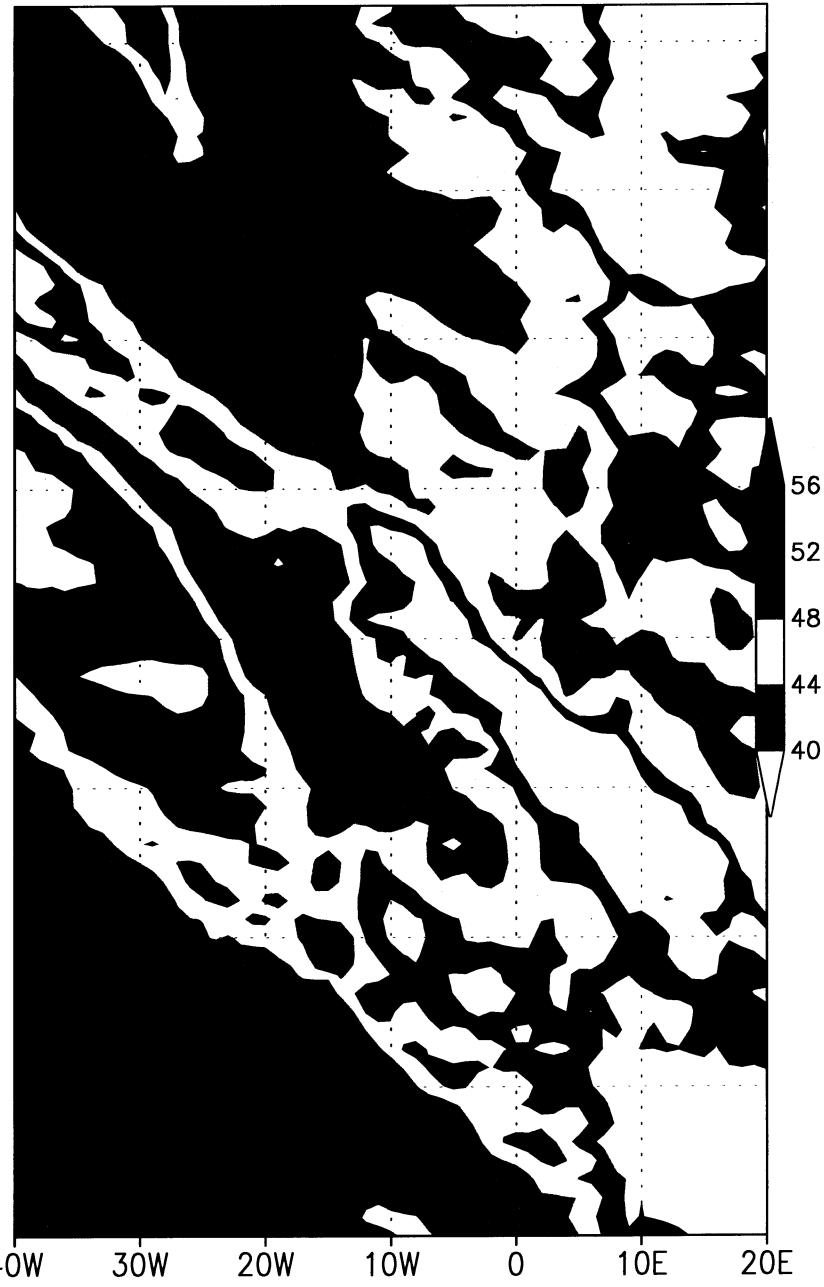
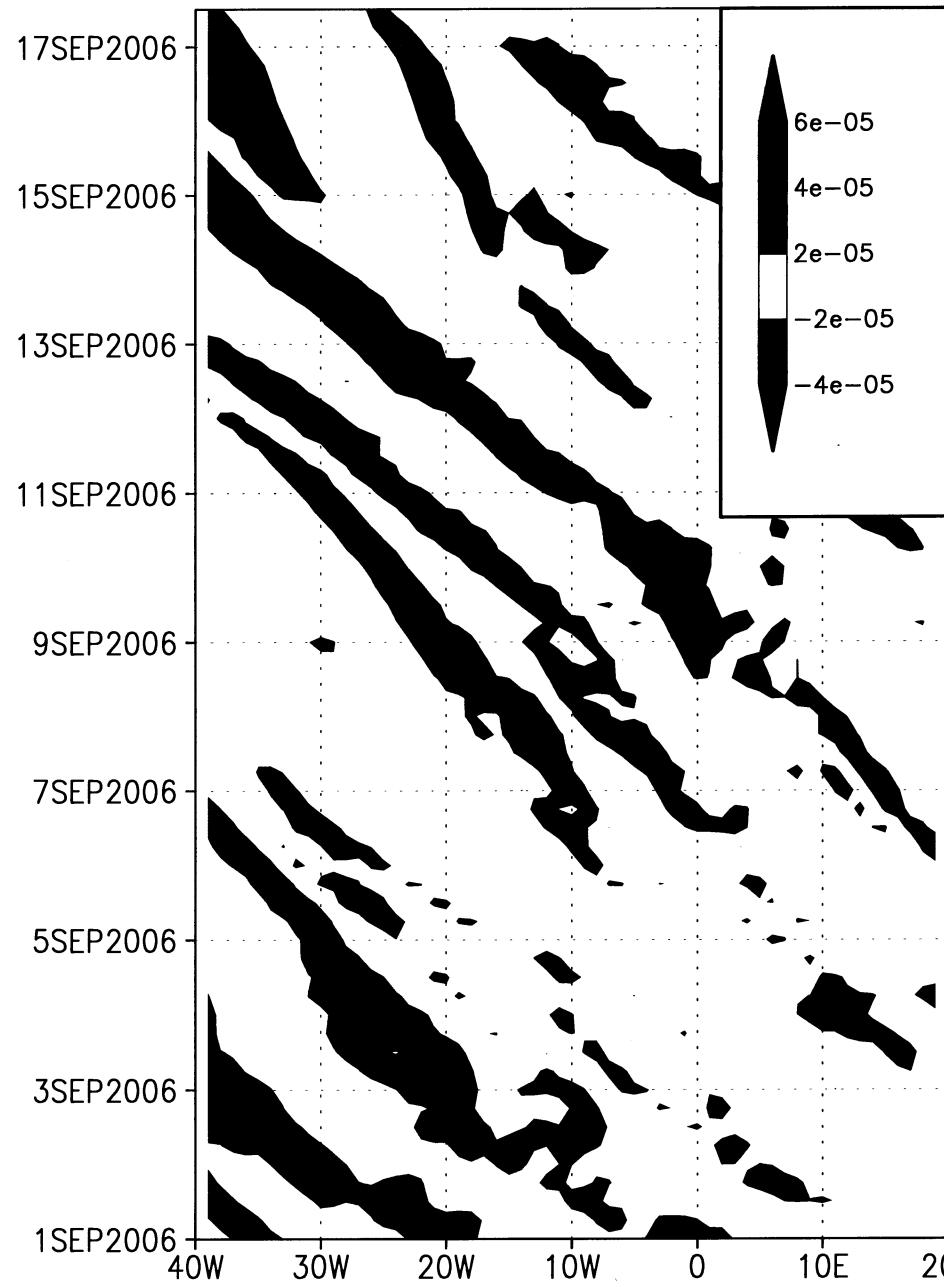
Data on pressure levels interpolated on a 1 degree resolution grid.

No significant difference from the original data in sigma levels (not shown).

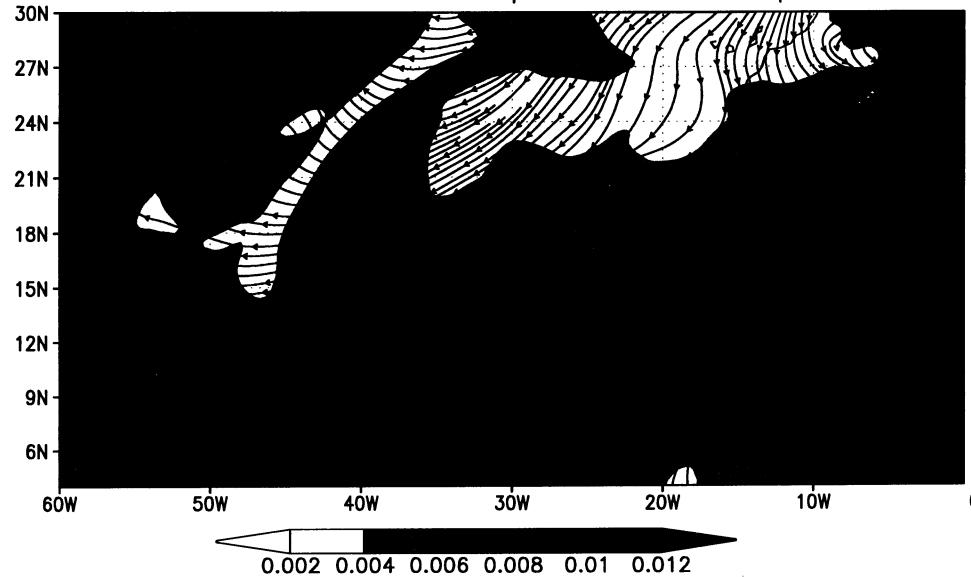
grl_fs02.eps

Caption:

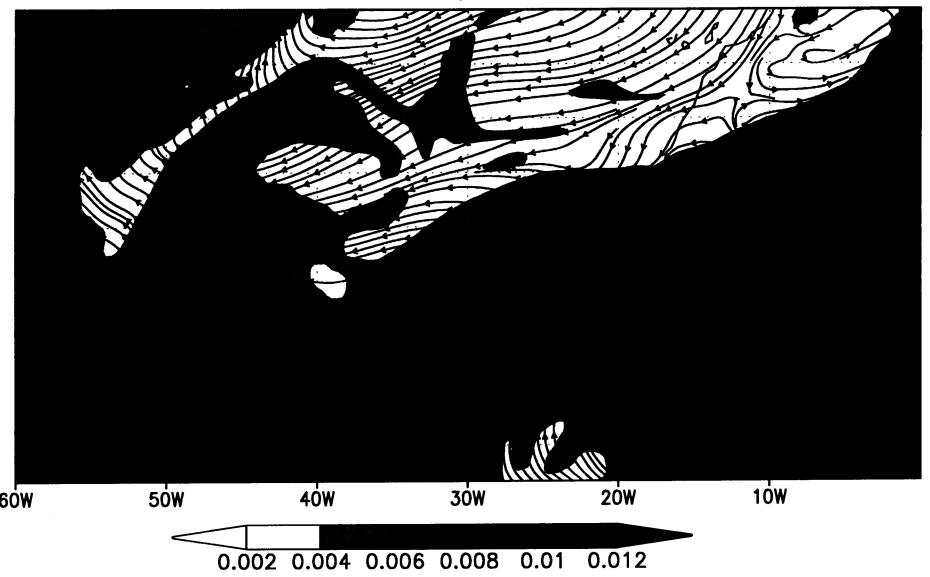
GEOS-5: 700 hPa specific humidity (Kg/Kg) and 850 hPa wind (streamlines) in the NCEP-derived initial conditions (upper left) for 00z 13 September, and relative to the 24, 48 and 72 hour forecasts for 00z 14, 15 and 16 September.



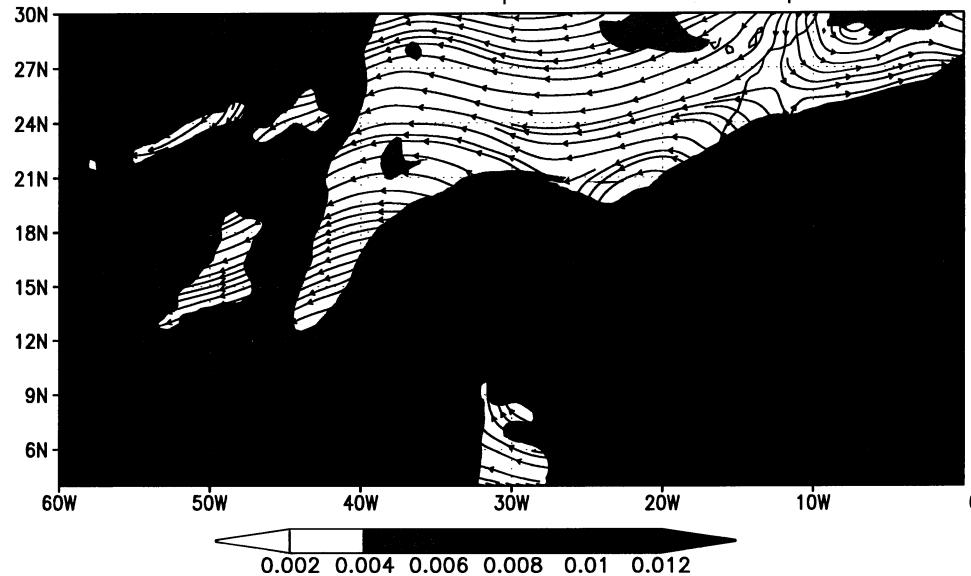
GEOS5 IN00z13Sep06 0hF VER13Sep



GEOS5 IN00z13Sep06 24hF VER14Sep



GEOS5 IN00z13Sep06 48hF VER15Sep



GEOS5 IN00z13Sep06 72hF VER16Sep

